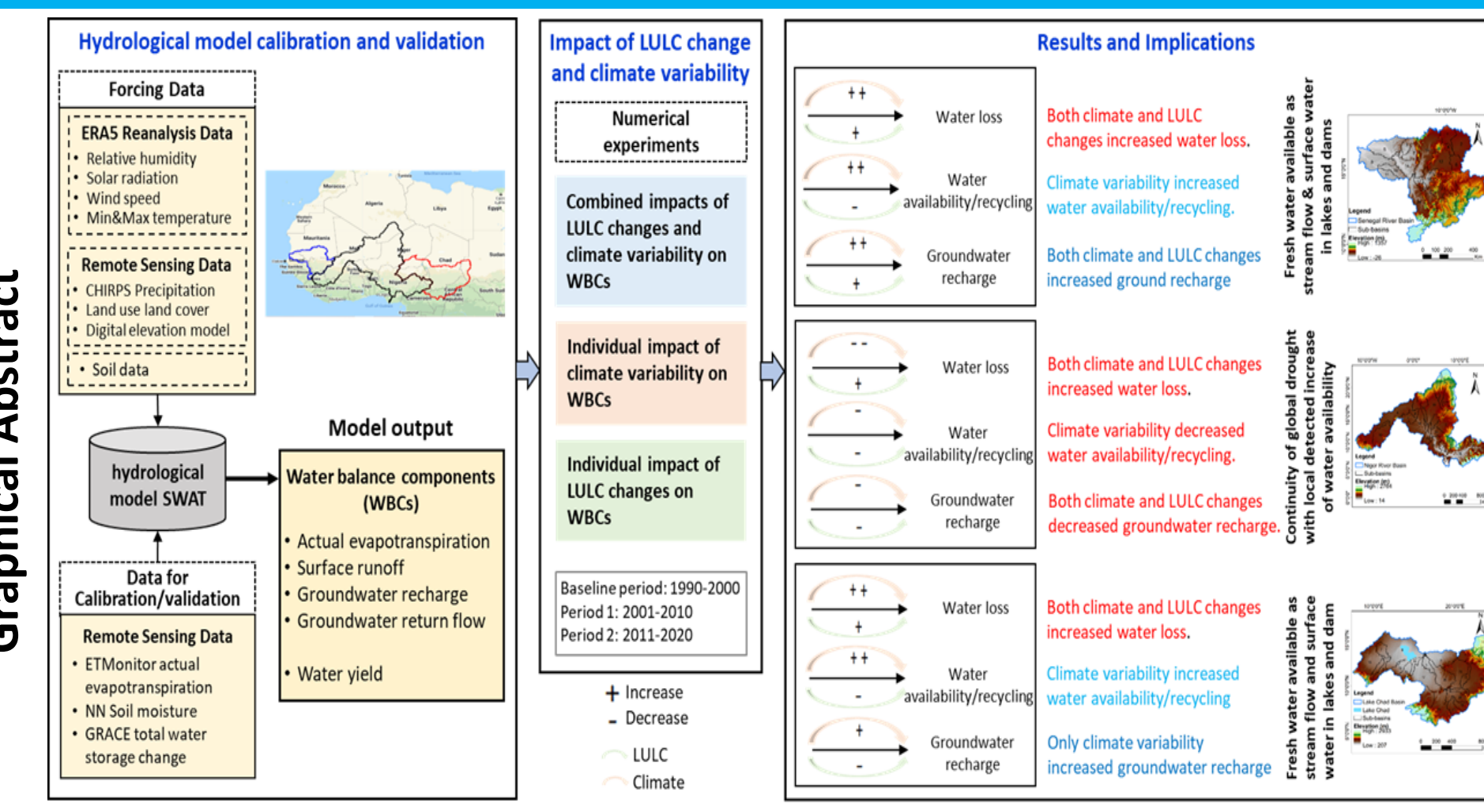


**Abstract:** The investigation of the hydrological responses to land use/land cover (LULC) change and climate variability is essential for understanding catchment hydrology, particularly in a vulnerable zone to global changes such as the Sahel region. Hence, our study contributed to separating and assessing the impacts of LULC change and climate variability on water balance components in the Sahel at the basin and sub-basin. Three basins were selected as study cases due to their importance in terms of catchment area (i.e. Senegal river, Niger river and Lake Chad basins). In this work, we have applied Soil and Water Assessment Tool (SWAT) model coupled with remote sensing retrievals of actual evapotranspiration (ETa) and surface soil moisture (SSM). To separate the impacts of the two aforementioned factors, two numerical experiments were designed: (i) climate variability effects by applying frozen LULC while changing the climate; (ii) LULC change impacts by applying frozen climate while changing LULC. The results revealed that, overall in the 2010s compared to the 1990s, the combined impact of LULC change and climate variability as well as separate effect of climate showed an increase in surface runoff, groundwater recharge and return flow in Senegal river and Lake Chad basins, while in Niger river basin most of all water balance components were declined. Frozen climate and change in LULC showed that spreading of natural vegetation at the expense of bare land led to an increase in actual ET and a decrease in surface runoff in the three watersheds, while in Senegal river basin it shows a slight increase in groundwater recharge and return flow. At sub-basin level, the analysis of LULC change showed that the gain in cropland and urban areas at the expense of the forest in some sub-basins, led to a local increase in surface runoff. This implies a better redistribution of water downstream and compensates the deficit in surface runoff caused by natural vegetation at the expense of bare land in some other catchments, i.e. a beneficial increase in fresh water availability. These changes at the same time with high intensity and long duration precipitation, this is likely to be a source of inundation and soil erosion in some small catchments in Niger river basin. Globally, the climate variability had a dominant impact on increasing water balance components resulting an increase in fresh water availability, with an extension and recovery of lake area in Lake Chad, which also increased groundwater return flow to rivers and water recycling within Senegal river and Lake Chad basins. In contrast, the LULC change was the major driver of decreasing the surface runoff, which could be a reason for lake area depletion in Lake Chad. At the same time, the two factors led to increasing water scarcity in Niger river basin. These outcomes emphasize the crucial role of water recycling which is the amount of water transferred from a sub-basin upstream to the next downstream within the watershed as well as give a good hydrological insight about water and land management in the study area. These findings are relevant to water resource management and to advance towards water-related Sustainable Development Goals (SDGs).

**Keywords:** African Sahel, SWAT model, ETMonitor, remote sensing soil moisture, LULC change, climate variability.



**Introduction:** Understanding the response of a watershed to climate variability is complex because it is non-linear due to the concurrent response of LULC to the same variability, response modulated by anthropic interventions. The non-linearity is due to the modification of a watershed hydrological properties because of LULC changes. This process is particularly relevant and complex under arid and semi-arid conditions such as in the Sahel, where changes in water availability lead to large and rapid changes in LULC, through the natural response of ecosystems to water and to efforts. A better understanding of the separate impacts of LULC and climate variability on water balance components is needed to develop effective land management policies towards sustainable water security.

Several studies paid attention to investigating the effects of LULC change and climate variability on hydrological response in the Sahel region using hydrological models or remote sensing data. Some studies have investigated the hydrological response to LULC change and climate variability in the Sahel before and during the drought [1]. Some studies have reported a recovery in vegetation and rainfall in the Sahel [2]. Some studies have investigated the hydrological response to LULC change and climate variability in the Sahel before the drought and in the post-drought period [3].

Almost of all previous studies were not able to separate the impacts of LULC change and climate variability due to the type of the used data (remote sensing data (NDVI) and ground observations data (precipitation and surface runoff)). Some other studies have applied hydrological models to separate the contribution of the two factors in changing hydrological response but in small catchments (<1000 km<sup>2</sup>). Most of the studies focused on surface runoff and streamflow, while other water balance components are equally important such as groundwater recharge and groundwater return flow. Moreover, water recycling as a consequence of LULC change was not evaluated. Most studies compared the drought period (1970s and 1980s) with the years past the drought till 2010. A study at basin level and sub-basin level, after the last recovery of vegetation and rainfall, has not been done yet.

**Impact score of LULC change and climate variability on water balance components**

To identify which factor dominated the impact, we calculated the impact scores of LULC change and climate variability following the

$$CR\_WBC_{CLI} = \Delta WBC_{CLI} / \Delta WBC$$

$$CR\_WBC_{LUI} = \Delta WBC_{LUI} / \Delta WBC$$

CR\_WBC\_CLI and CR\_WBC\_LUI are the separate impact scores on water balance components, i.e., actual ET, SW, GW\_RCH, and SURQ, caused by climate variability and caused by LULC change, respectively; ΔWBC is the change in a water balance component caused by the combined impact of LULC change and climate variability relative to baseline.

**Performance metrics used to evaluate the calibration and validation [7]**

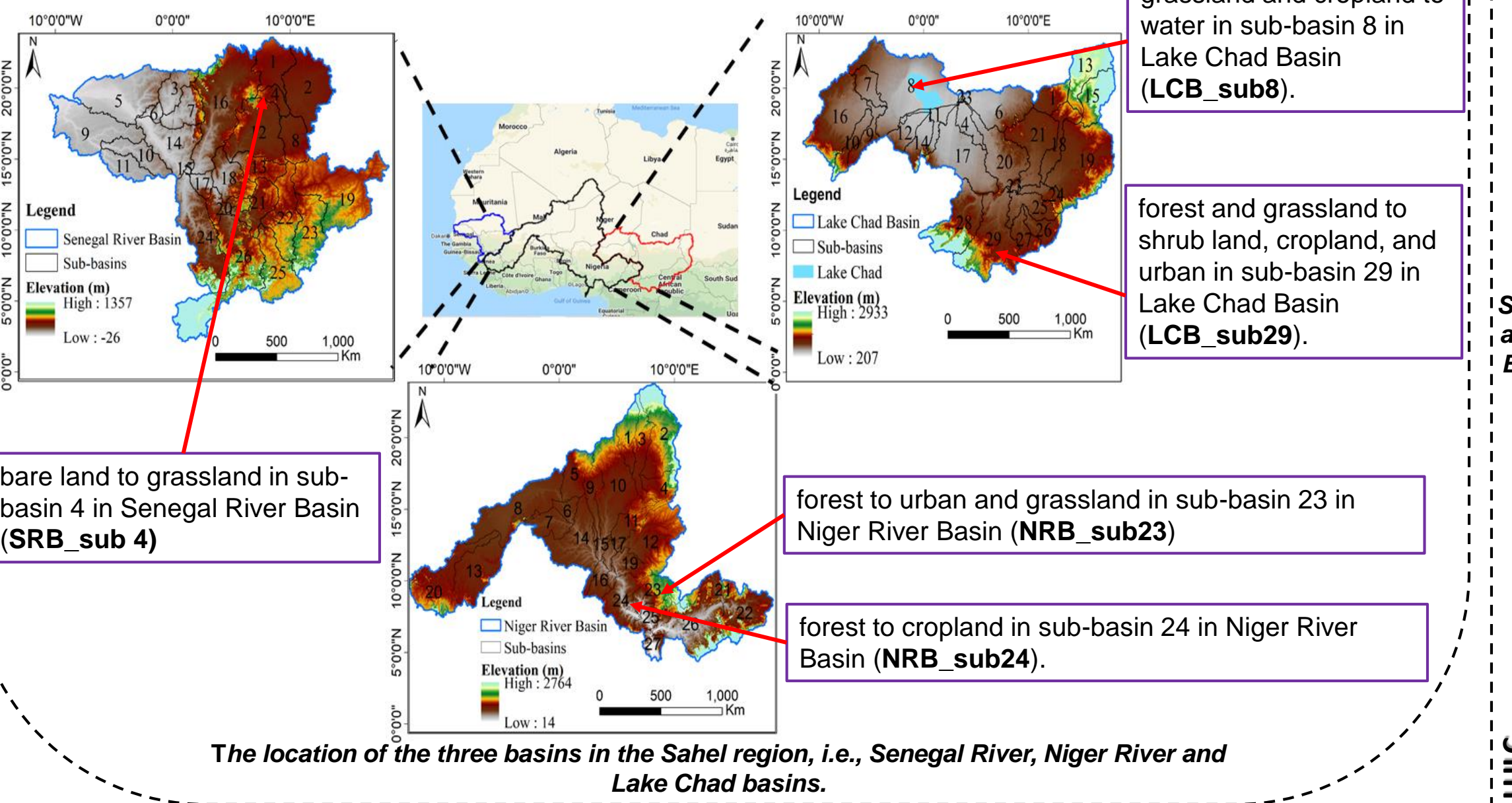
Performance Metrics	Equations	Descriptions
Coefficient of determination	$R^2 = \frac{(\sum (ET_{sim} - \bar{ET}_{sim})(ET_{sat} - \bar{ET}_{sat}))^2}{\sum (ET_{sim} - \bar{ET}_{sim})^2 \sum (ET_{sat} - \bar{ET}_{sat})^2}$	where $ET_{sat}$ represents satellite-based ETa values; $ET_{sim}$ represents simulated ETa values; $\bar{ET}_{sat}$ represents mean satellite-based ETa values; $\bar{ET}_{sim}$ represents mean simulated ETa values.
Nash-Sutcliffe Efficiency	$NSE = 1 - \frac{\sum (ET_{sim} - ET_{sat})^2}{\sum (ET_{sat} - \bar{ET}_{sat})^2}$	$r$ is the Pearson product correlation coefficient between satellite-based ETa and the simulated ETa; $\alpha$ is the standard deviation of the simulated ETa over the standard deviation of the satellite-based ETa; $\beta$ is the ratio of the mean simulated ETa to the satellite-based ETa.
Kling-Gupta Efficiency	$KGE = 1 - \sqrt{(r-1)^2 + (\alpha-1)^2 + (\beta-1)^2}$	
Percent bias	$PBIAS = \frac{\sum (ET_{sim} - ET_{sat})}{\sum ET_{sat}}$	

**Objective:**

- 1) to evaluate the separate impacts of climate variability and LULC on water balance components using a physically-based hydrological model
- 2) to estimate the responses of different water balance components to climate variability and LULC changes at basin and sub-basin scale
- 3) to compare the relative contributions of these different factors and identify which one has the dominant impacts

**Study area**

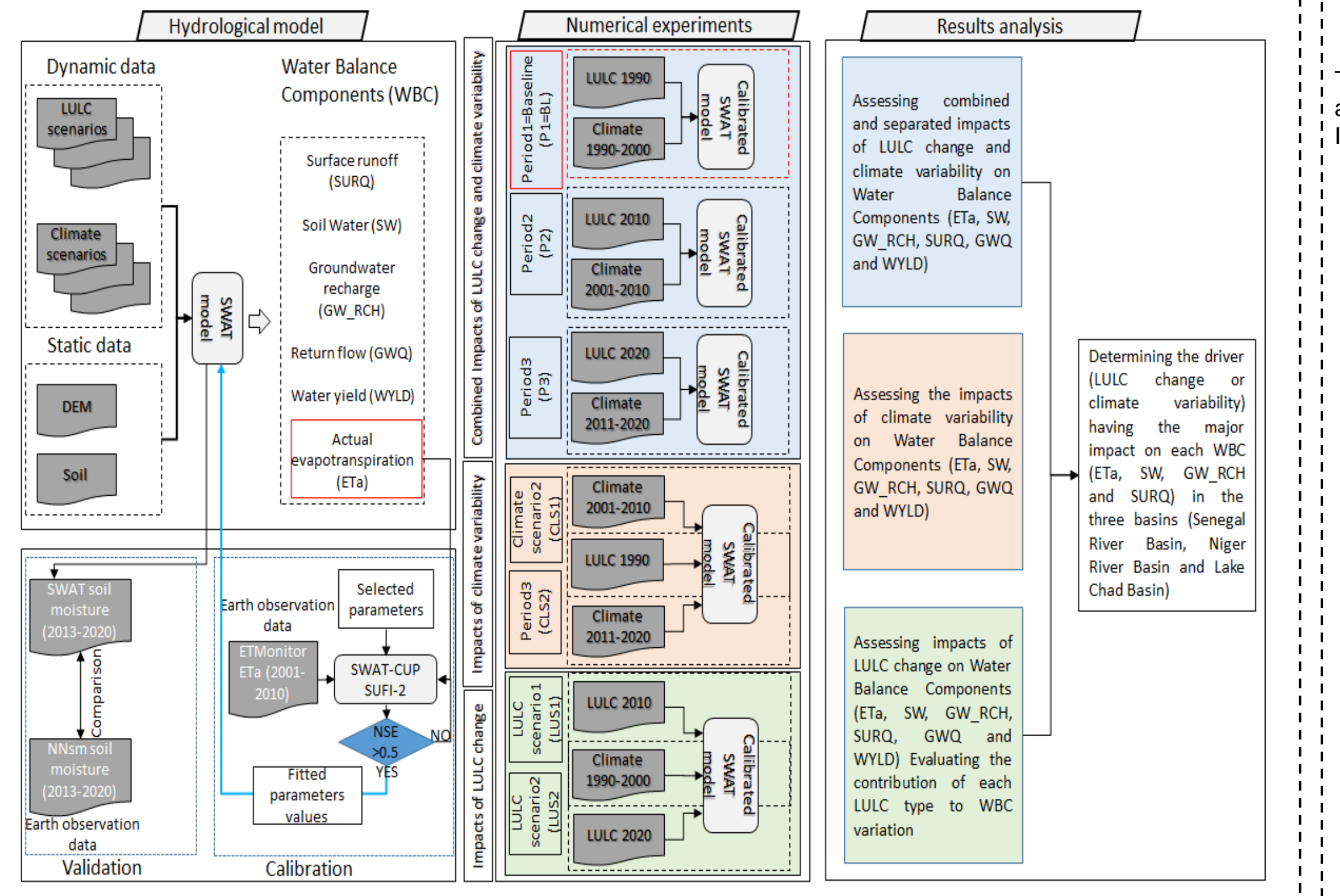
The Senegal River basin (SRB) is located in the western part of the Sahel region, with a total drainage area of approximately 375,000 Km<sup>2</sup> and a length of 1800 km. The region is inhabited by 3.5 million people. The Niger River basin (NRB) has a catchment area of 2.1x10<sup>6</sup> km<sup>2</sup> and a length of 4200 km. This region is inhabited by more than 100 million people and spans nine countries. The Lake Chad Basin (LCB) is the largest endorheic lake basin in the world, and it is located in the center of the African Sahel between 5.19–25.29°N latitude and 6.85–24.45°E longitude with an area of 2.5 10<sup>6</sup> km<sup>2</sup>. The region is inhabited by 17.4 million people.



**Method**

Hydrological model: Soil and Water Assessment Tool (SWAT) [4] which was calibrated and validated by using remote sensing data, and this was well described in our previous paper [5]. Then the calibrated SWAT was used to separate and assess the impacts of LULC change and climate variability on WBCs. The study period was split into 3 sub periods:

Period 1 =Base Line (P1=BL)= Climate 1990–2000 with LULC1990, Period 2 (P2)= Climate 2001–2010 with LULC2010 and Period 3 (P3)= Climate 2011–2020 with LULC2020. Three climate scenarios were used in this study, Climate 1990–2000=base line (BL), Climate 2001–2010=Climate Scenario 1 (CLS1) and Climate 2011–2020=Climate Scenario 2 (CLS2). Three LULC scenarios were used in this work, LULC1990=base line (BL), LULC2010=Land Use Scenario 1 (LUS1) and LULC2020= Land Use Scenario 2 (LUS2).



**Study workflow:** 1) Hydrological modeling including Calibration/Validation of SWAT model; 2) Numerical Experiments including assessment of the combined impact of LULC and climate change, the impact only by climate change and the impact only by LULC change on each of the simulates water balance components; 3) Results Analysis

**(1) The combined impacts by LULC change and climate variability on WBCs**  
It is calculated following the following equation:  $\Delta WBC_{P_i} = WBC_{P_i} - WBC_{BL}$   
ΔWBC<sub>P<sub>i</sub></sub> is the change in a water balance component between the corresponding period (P<sub>i</sub> = P<sub>2</sub> or P<sub>3</sub>) and the baseline period (P<sub>1</sub>). WBC<sub>P<sub>i</sub></sub> is the value of the water balance component in the corresponding period, WBC<sub>BL</sub> is the value of a water balance component in the baseline period

**(1) The impacts of climate variability on WBCs**  
It is calculated following the following equation:  $\Delta WBC_{CLI} = WBC_{CLS_i} - WBC_{BL}$   
ΔWBC<sub>CLI</sub> is the change in a water balance component between the corresponding period (CLS1 or CLS2) and the baseline period, WBC<sub>CLS<sub>i</sub></sub> is the value of the water balance component for either CLS1 or CLS2, WBC<sub>BL</sub> is the value of a water balance component in the baseline period.

**(1) The impacts of LULC changes on WBCs**  
It is calculated following the following equation:  $\Delta WBC_{LUI} = WBC_{LUS_i} - WBC_{BL}$   
ΔWBC<sub>LUI</sub> is the change in a water balance component between the corresponding period (LUS1 or LUS2) and the baseline period, WBC<sub>LUS<sub>i</sub></sub> is the value of a water balance component for either LUS1 or LUS2, WBC<sub>BL</sub> is the value of a water balance component in the baseline period.

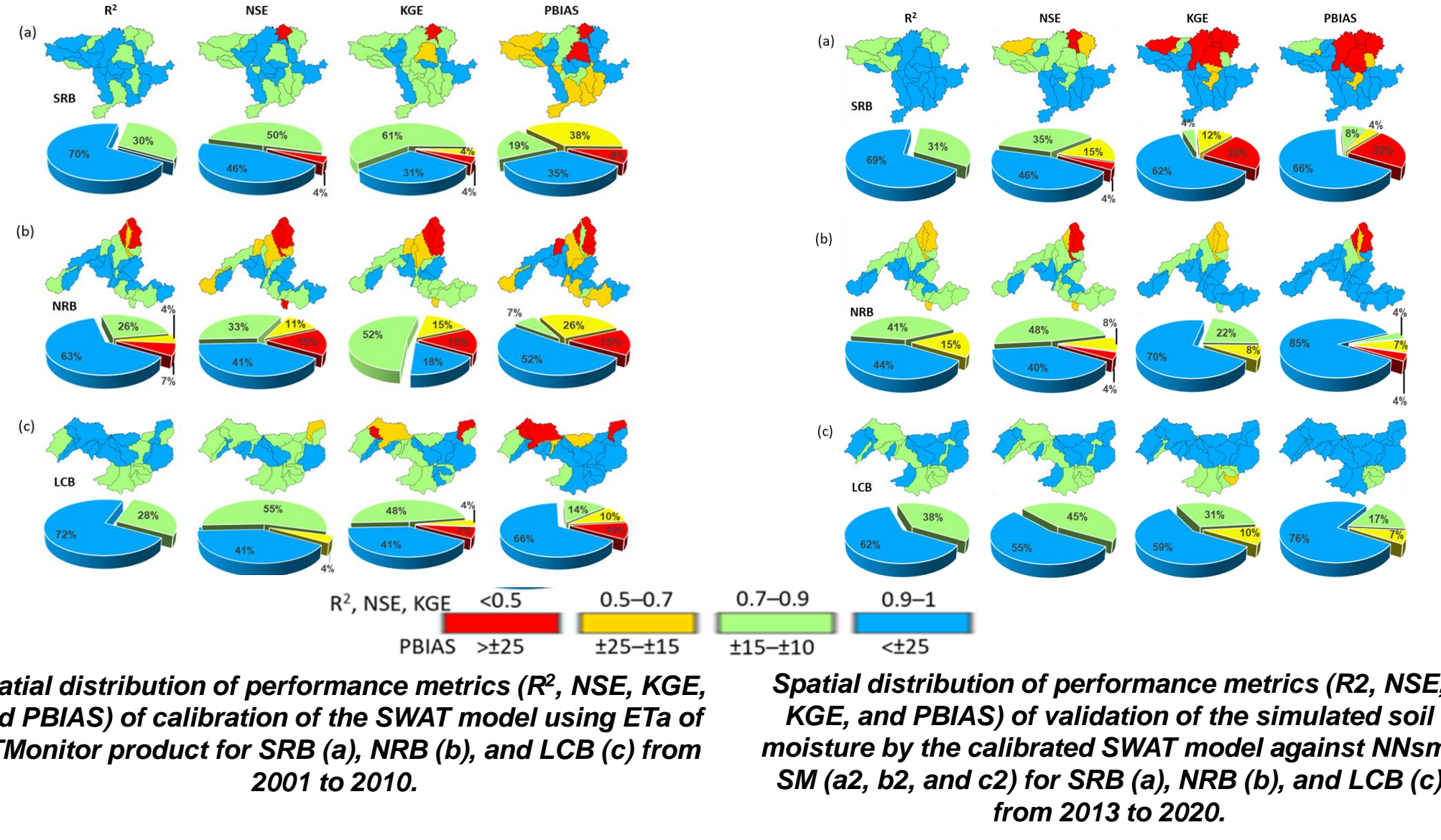
**An indicator of water recycling: the ratio of Water outflow to Water inflow**

The SWAT provides options to describe the use of streamflow in the river channels and allows inter-sub-basin water transfers [6]. We proposed to use the ratio of WYLD/PREC as an indicator for potential water recycling, i.e. the ratio of out- recycling (R<sub>out</sub>), calculated as:  
 $R_{out} = WYLD / PREC$   
where R<sub>out</sub> is the ratio of out recycling, PREC is the sum of precipitation over all sub-basins (km<sup>3</sup>), WYLD (km<sup>3</sup>) is the net amount of water that leaves the sub-basin and contributes to streamflow in the reach and calculated as:  
 $WYLD = SURQ + LATQ + GWQ - Q_{LOSS}$   
where LATQ (km<sup>3</sup>) is the lateral flow contribution to streamflow; and Q<sub>LOSS</sub> (km<sup>3</sup>) is the transmission loss.

**Results**

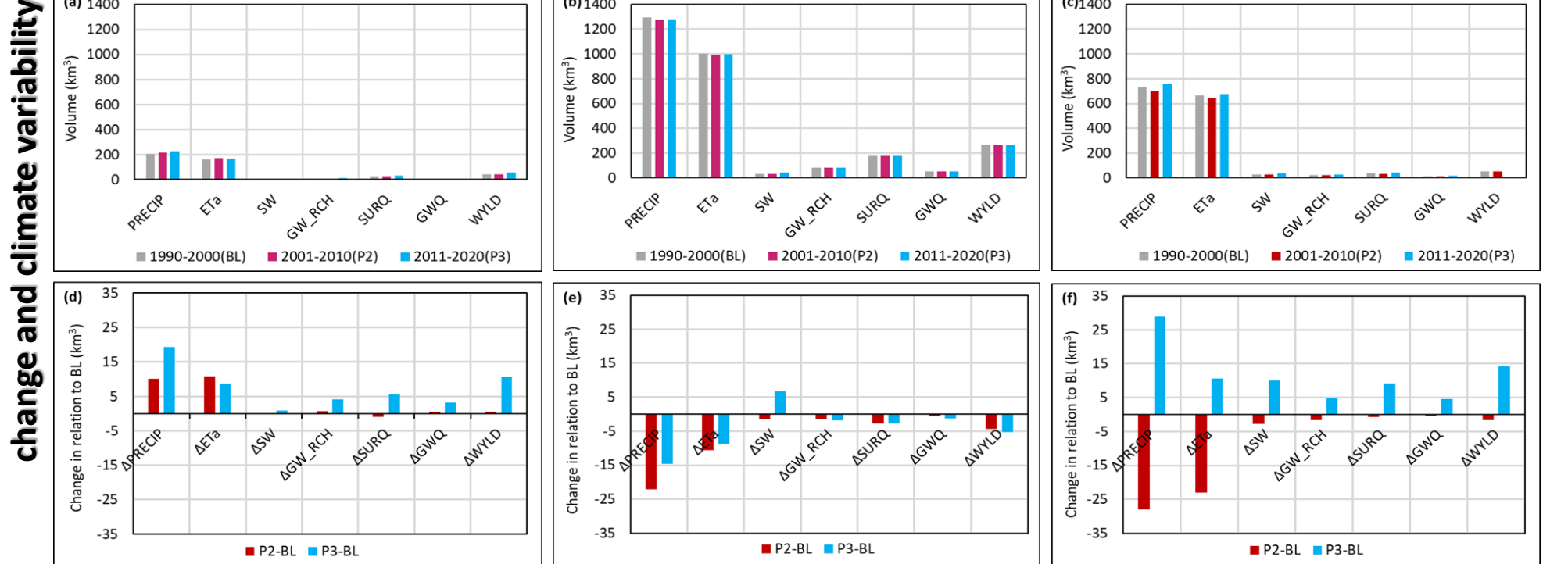
**Calibration based on actual evapotranspiration and validation by using soil moisture**

In the three basins, the calibration using monthly ETMonitor retrievals of actual ET indicated a good performance with values of R<sup>2</sup>, NSE, and KGE greater than 0.7 and PBIAS values lower than ±15%. The validation for the three basins using monthly NNsm soil moisture showed that R<sup>2</sup>, NSE, and KGE were higher than 0.7 and PBIAS lower than ±15%.



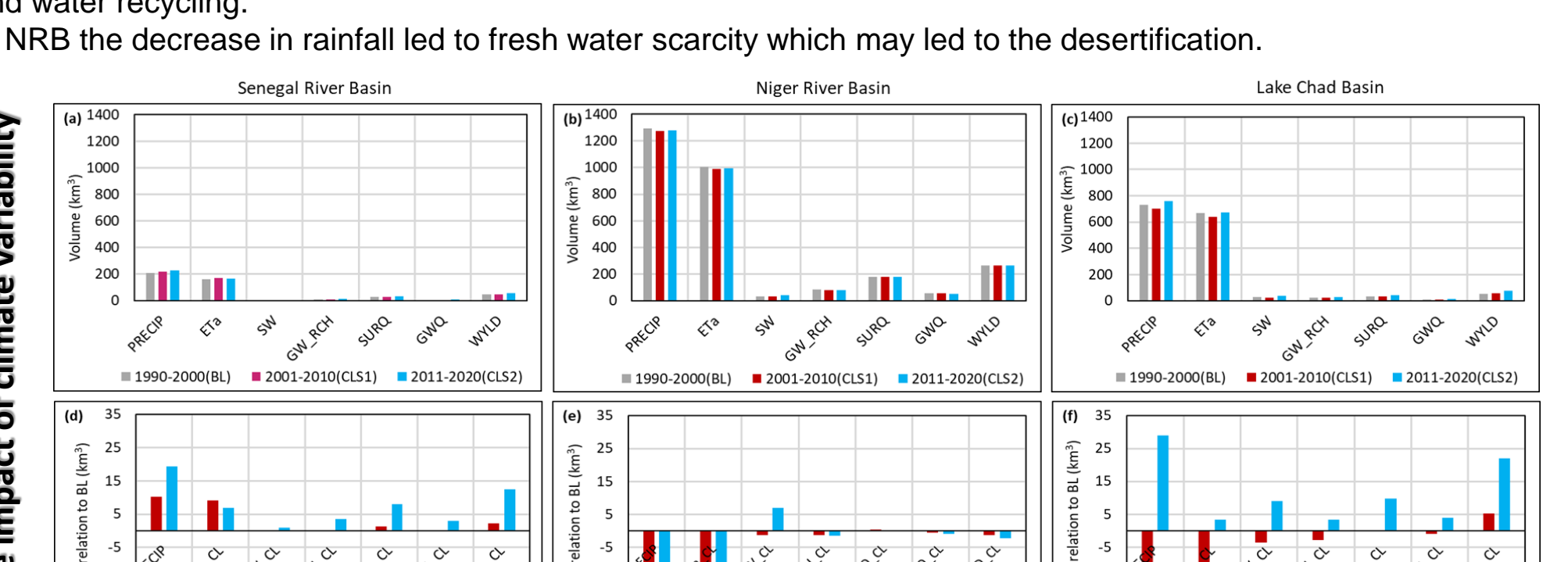
**Combined Impacts of LULC changes and climate variability on water balance components**

The change in WBCs were evaluated during P2 (2001–2010) and P3 (2011–2020) compared to P1=BL (1990–2000). In P2 (2001–2010), all WBCs decreased in NRB and LCB, while in contrary in SRB most WBCs increased except for surface runoff. In P3 (2011–2020), all WBCs increased in SRB and LCB, while in NRB most WBCs decreased except for soil water content (SW).



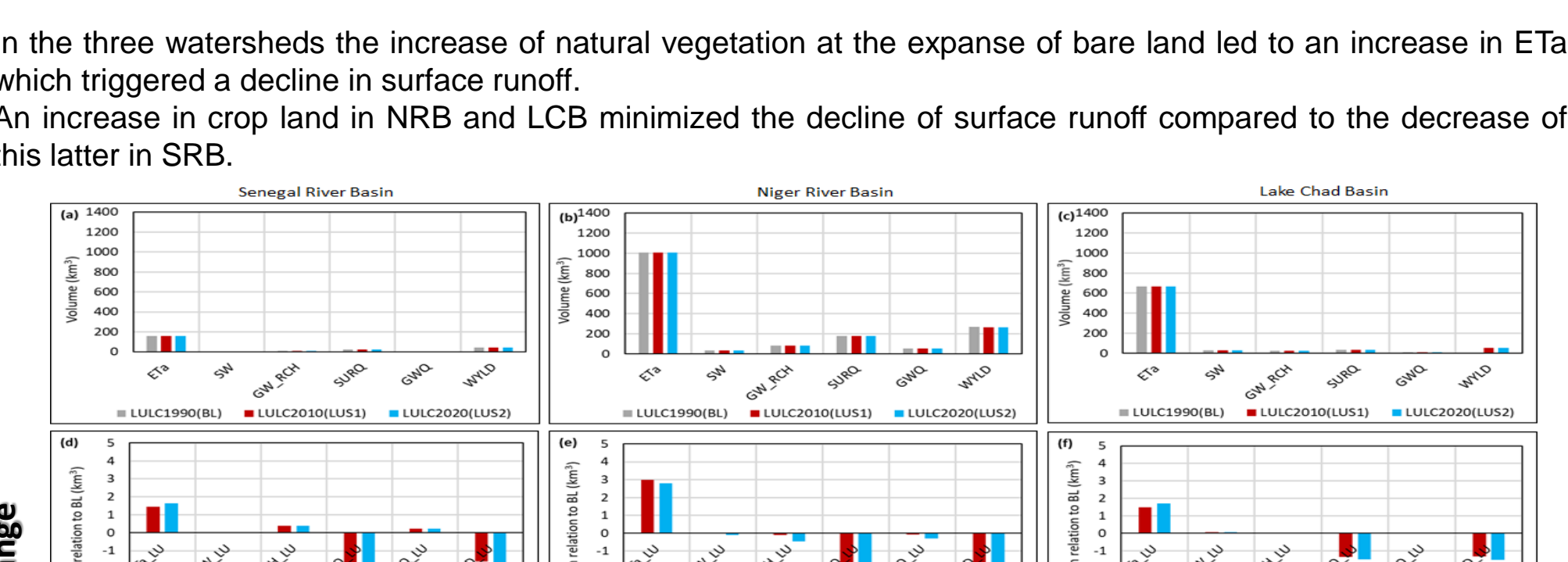
**separate impacts of climate variability on water balance components**

The increase in rainfall in SRB and LCB led to an increase in most of all WBCs specifically fresh water availability and water recycling. In NRB the decrease in rainfall led to fresh water scarcity which may led to the desertification.



**separate impacts of LULC changes on water balance components**

In the three watersheds the increase of natural vegetation at the expense of bare land led to an increase in ETa which triggered a decline in surface runoff. An increase in crop land in NRB and LCB minimized the decline of surface runoff compared to the decrease of this latter in SRB.



**Discussion**

- In the SRB, the rise in SURQ, GW\_RCH, and return flow (GWQ), during both CLS1 (2001–2010) and CLS2 (2011–2020) compared to baseline (1990–2000), [Oyebande and Oduunuga \(2010\)](#) [8].
- In the NRB, the decline in SURQ, GW\_RCH, and GWQ caused by the deficit in rainfall during both CLS1 and [Descroix et al. \(2018\)](#) [9].
- In LCB, the decline in SURQ during CLS1 was clearly caused by rainfall deficit, [Mahmood and Jia \(2019\)](#).
- The re-greening of the region by expansion of grassland and shrub land led to an increase in actual ET and a decrease in SURQ, [Yonaba et al. \(2021\)](#) and [Ogutu et al. \(2021\)](#). This vegetation recovery played a major role in increasing GW\_RCH and soil water content (SW), [Marshall et al. \(2012\)](#). This finding applied to three watersheds, but the change was larger in NRB and LCB than that in SRB, [Mahmood and Jia \(2019\)](#).
- Both climate variability and LULC change had nearly similar impacts on the hydrological response in the SRB. The LCB was more sensitive to the LULC change during P2. This process significantly reduced water availability by decreasing the SURQ and increasing the water loss via evapotranspiration, because of the higher actual ET. In contrast, climate variability was the main driver of increasing SURQ during P3, [Pham-Duc et al. \(2020\)](#) [10].

**Conclusions**

- An increase in rainfall in both the SRB and the LCB, leading to a rise in freshwater availability and water recycling because of the increase in water yield. In the NRB, the shortage of water supply continued, resulting in the continuation of the drought in the basin, which showed that the Sahelian hydrological paradoxes did not apply at basin level.
- In LCB and SRB, the increase in water yield, caused by climate variability, played an important role in improving water availability in the basin.
- The man-made extension of cropland at the expense of the forest loss had an important role in mitigating the decrease in surface runoff.
- The continuous population growth led to an increase in the cultivated area at the expense of forest loss, i.e., a further increase in surface runoff. The latter may contribute to water recycling through water transfer from a sub-basin upstream to the next sub-basin downstream. These findings emphasized the crucial role of water recycling within the watershed, as well as giving a good hydrological insight into the interrelation of water and land management in the study.

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